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HIGGINSITE, A NEW MINERAL OF THE OLIVENITE GROUP

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The new mineral described in this paper was found by J. B. Tenney at the Higgins Mine, Bisbee, Arizona and was sent to the Harvard Mineralogical Museum for determination in 1917. Mr. Tenney, then Chief Geologist of the Phelps-Dodge Corporation, described the occurrence of the mineral in the following words:—"The mineral occurs always with manganese ore on or near the surface. The manganese minerals are psilomelane, braunite and pyrolusite. Occasionally barite occurs with it also, and the manganese passes on the edge of the deposits into limonite (or goëthite) intergrown with cherty quartz, or into crystalline limestone. The form and nature of the occurrence suggest replacement of pure limestone by manganese, iron and silica, carried in strong sulfate solutions. The green mineral appears to be of the same age, as far as can be told."

The authors take pleasure in following the suggestion of Mr. Tenney that this mineral bear the name *higginsite*, after the Higgins Mine, where it was found.

The specimens received show crystals and granular masses of higginsite interspersed thru black manganese ores with an occasional plate of white barite. The color of the crystals is a vivid malachite green; of the granular material, yellow green. The powder is distinctly yellowish green, with a brownish cast. In the coarsest material grains as much as 2 cm. in diameter were seen, but the largest crystal measured was about 5 mm. in length.

Crystallography—Higginsite is orthorhombic, with the elements:

$$p_0 = 1.272 \quad q_0 = .7940.$$

Details are given on a later page. Its close relationship in

form to olivenite and descloizite is shown by comparison of their axial ratios:

	a	b	c
Higginsite	.6242	1	.7940
Descloizite	.6368	1	.8045
Olivénite ($\frac{2}{3}a$)	.6264	1	.6726

Figures 1 and 2 show the prevailing habits of the crystals. The prismatic habit is most frequent in those at hand; but the domatic development, either as shown or with nearly equal extension parallel to the two horizontal axes, gives some crystals an octahedroid appearance. The lack of any pronounced cleavage makes the orientation of the crystals difficult. The faces, especially of the pyramid forms, are of good luster and gave on the whole very satisfactory reflections on the goniometer.

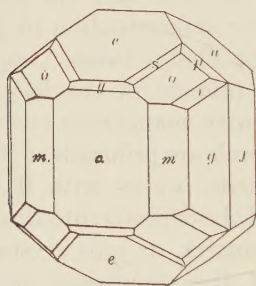


FIG. 1

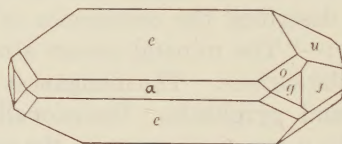


FIG. 2

Physical Characters.—Hardness about 4.5; specific gravity 4.33. The optical properties were studied by Mr. T. Matsumoto and partly determined as follows, the small size and opacity of the crystals rendering complete results unattainable:

Refractive index $\beta > \alpha > 1.745$; birefringence approx. 0.030.

Optical orientation: $X \parallel a$, probably Bx_a ; $Y \parallel b$; $Z \parallel c$; Axial plane therefore parallel to (010).

Opt. character:—(?); axial angle large; dispersion $v > \rho$ (if —).

Absorption $Y > X > Z$; pleochroism marked: X green; Y yellow green; Z blue green.

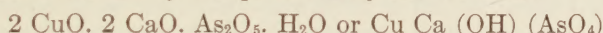
Chemical Composition.—The analyses (by E. V. S.) were made on carefully hand-picked grains, freed as far as possible from manganiferous gangue. Some of the latter appeared to be present as dust even in the clearest crystals and could not be wholly removed. The final analysis was made upon about 1 gram of material, the check determination of copper and calcium on a half gram of the same powder.

TABLE 1
ANALYSES OF HIGGINSITE, BY E. V. SHANNON

	1	2	Av.	Ratios		
CuO.....	28.51	28.82	28.67	0.3603	1.93	2
CaO.....	20.92	20.74	20.83	0.3712	1.99	2
V ₂ O ₅	1.97		1.97	} 0.1901	1.02	1
As ₂ O ₅	41.23		41.23			
H ₂ O - 105° C.....	0.08		0.08	} 0.1892	1.02	1
H ₂ O + 105° C.....	3.41		3.41			
Fe ₂ O ₃	0.48		0.48			
MnO.....	2.84		2.84			
Gangue (Insoluble)...	0.86		0.86			
Total.....	100.30		100.37			

All the iron and manganese is regarded as extraneous and derived from the gangue material. Deducting these, with the insoluble material and hygroscopic water, the average analytical figures yield the ratios in the last column.

These ratios then yield quite closely the formula:



Arsenic is replaced to a small extent by vanadium. This may be compared with the formulas of olivenite and descloizite: 4 CuO. As₂O₅. H₂O, and 2PbO. 2 ZnO. V₂O₅. H₂O, which are typical of the olivenite group.

Recalculating the essential constituents of the mineral to 100% and comparing with these figures the composition required by the formula derived above we obtain the following figures.

	Original	Recalculated to 100 %	Theory for Cu Ca OHAsO ₄
CuO.....	28.67	29.83	30.65
CaO.....	20.83	21.67	21.60
V ₂ O ₅	1.97	2.05	} 44.28
As ₂ O ₅	41.23	42.90	
H ₂ O.....	3.41	3.55	3.47

These figures show a very satisfactory agreement.

Pyrognostic Characters.—Higginsite fuses at about 3, coloring the flame at first pale blue (As) and then blue green (Cu). On charcoal alone it fuses to a black slag without flame coloration or coating; with soda and borax it is reduced to metallic copper, yielding a faint arsenic reaction. In the closed tube it decrepitates slightly, turns black, and at a red heat gives off a little neutral water. Gives no arsenic sublimate when heated with charcoal in the closed tube. It is readily soluble in nitric and hydrochloric acids; partially soluble in sulfuric acid; insoluble in ammonia.

THE GOLDSCHMIDT TWO-CIRCLE METHOD. CALCULATIONS IN THE ORTHORHOMBIC SYSTEM

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The following relations may be derived from the diagram, figure 32, in which pq is the face-pole of a pyramid for which the angles φ and ρ are known:

$$x = \sin \varphi \tan \rho = pp_0; y = \cos \varphi \tan \rho = qq_0$$

From these, p and q being known, p_0 and q_0 can be calculated.

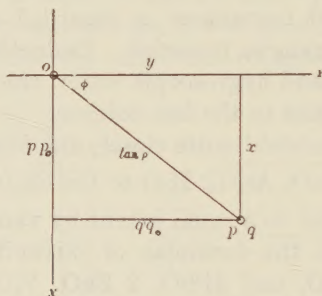


FIG. 32

Each face gives values for x and y and the averages for n faces have the form:

$$p_0 = \frac{1}{n} \left(\frac{x_1}{p_1} + \frac{x_2}{p_2} + \dots + \frac{x_n}{p_n} \right)$$

$$q_0 = \frac{1}{n} \left(\frac{y_1}{q_1} + \frac{y_2}{q_2} + \dots + \frac{y_n}{q_n} \right)$$

For domes: If $y = 0$, $x = pp_0$; If $x = 0$, $y = qq_0$.

For prisms:

$$\text{If the angle } hk0 \text{ to } 010 = \varphi, \tan \varphi = \frac{pp_0}{qq_0}.$$

RELATIONS OF ELEMENTS TO LINEAR AXES

$$p_0 = \frac{c}{a}. \quad q_0 = c. \quad a = \frac{q_0}{p_0}. \quad c = q_0$$

CALCULATION OF ANGLES FROM ELEMENTS

From the diagram and equations of the first paragraph we have for any face, pq :

$$\frac{x}{y} = \frac{pp_0}{qq_0} = \tan \varphi; \frac{x}{\sin \varphi} = \frac{y}{\cos \varphi} = \tan \rho; \sqrt{x^2 + y^2} = \tan \rho$$

For domes: If $x = 0$, $\tan \varphi = 0$, $\varphi = 0$; $\tan \rho = qq_0$.

If $y = 0$, $\tan \varphi = \infty$, $\varphi = 90^\circ$; $\tan \rho = pp_0$.

For prisms:

$$\frac{p}{q} \infty, \tan \varphi = \frac{pp_0}{qq_0}; \tan \rho = \infty, \rho = 90^\circ.$$

ILLUSTRATION OF THE ORTHORHOMBIC SYSTEM. MEASUREMENTS AND CALCULATIONS ON HIGGINSITE

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As an illustration of the application of the formulas given in the preceding article, the following discussion of a crystal of the new mineral higginsite, described above, may well serve. The measurements of one crystal are given in full in Table 1, the form of calculation being that used thruout in Goldschmidt's work. In this table, columns 1, 2, 5, and 6 contain the record of actual observation on the goniometer. The numbers of col. 1 are those used to mark the faces in the note-book sketch of the crystal; the letters of col. 2 stand for good, fair and poor, depending on the quality of the reflected signals; col. 5 contains the angles read on the vertical circle, V, col. 6 those on the horizontal circle, H, of the goniometer.

These angles were plotted in gnomonic projection yielding a diagram similar to figure 33. The next step was the choice of the unit form. Either of the pyramids, o and p , might have been taken for this and its coördinates would then have been the elements, p_0 and q_0 . The choice fell upon o because this form is more prominently developed on the crystals; the zonal relations with other forms are at least as good; and its selection brings to expression the isomorphism of the new species with descloizite, as will be shown below.

The unit form chosen, the Goldschmidt symbols could be read at once from the projection; they are entered in col. 3. The letters of col. 4 follow the usage for the mineral descloizite.

Determination of the value v_0 was next in order. The projection showed that the face 1 will have $\varphi = 0$, and therefore v_0 would be close to $77^\circ 28'$, the V reading of face 1. Each of the pairs of faces: 2 and 3; 4 and 5; 8 and 9; are symmetrically dis-

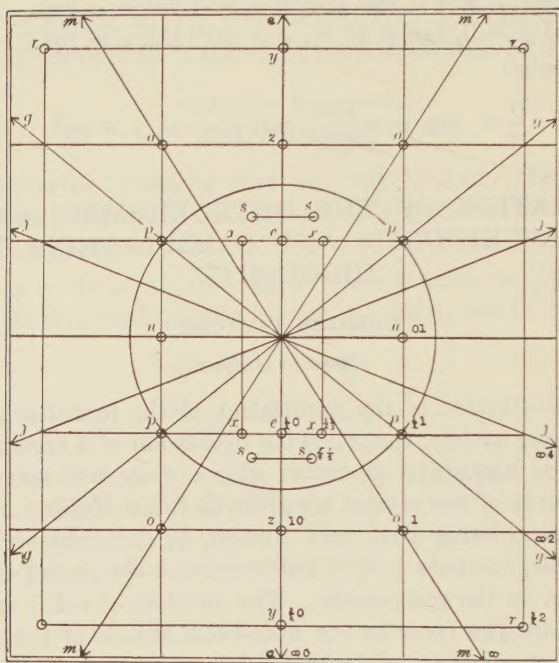


FIG. 33

posed to 1, and the half-sums of the V readings of each pair gave independent values for v_0 . Faces 6 and 7 should be 90° from face 1 in φ , and the prism faces also yielded, either directly or by taking the half-sum of symmetrical readings, other values of v_0 . The average of all is $77^\circ 29'$ which, subtracted from each angle in col. 5, gives the values of φ entered in col. 7.

Col. 8 contains the values of ρ obtained from col. 6 by subtracting each from $h_0 = 260^\circ$, a constant for the instrument.

The calculation of p_0 and q_0 followed. In col. 9 was written first, $\lg \tan \rho$ for each face (except prisms); then $\lg \sin \varphi$ and $\lg \cos \varphi$ respectively above and below the first. Col. 10 contains the sums of these logarithms, above the upper two, below the lower two, the logarithms respectively of x and y . This addition proceeds most rapidly if done beginning at the left hand side of the numbers to be added, the sum being then written in col. 10 in order from left to right, a trick of addition (or subtraction)

TABLE 1, on following page and half of page 162, gives the measurements on the higginsite and the calculations therefrom.

1	2	3	4	5	6	7	8	9	10	11	12	13
No.	Qual.	Symbol	Letter	v	h	$\varphi_v = 77^\circ 29'$	$\rho_{ho} = 260^\circ$	$\lg \sin \varphi$ $\lg \tan \rho$ $\lg \cos \varphi$	$\lg x$ $\lg y$	x = pDo y = qDo	dPo	qo
1	G	01	u	77°28'	221°46'	00°01'	38°14'	∞ 989645 0	989645	.7879		.7879
2	G	½1	p	38 44	214 37	-38 45	45 23	979652 000581 989203	980233	.6343	.6343	.7904
3	P	½1	p	116 15	214 37	38 46	45 23	979668 000581 989193	980249	.6346	.6346	.7904
4	P	1	o	19 35	203 49	-57 54	56 11	992795 017401 972542	989774	.7902	.6320	.7902
5	F	1	o	135 15	204 06	57 46	55 54	992731 016938 972703	989943	.7933	.6245	.7933
6	P	½0	e	348 25	227 37	-89 04	32 23	999994 980223 ∞	989641	.7878	.6341	.7878
7	G	½0	e	167 00	227 33	89 31	32 27	999998 980335 ∞	980333	.6358	.6358	
8	F	¾2	r	127 19	191 56	49 50	68 04	988319 039505 980957	027824	1.898	.6326	.8010
9	F	¾2	r	27 13	191 58	-50 16	68 02	988594 039432 980565	028026	1.907	.6356	.7925
10	P	¾¼	s	153 55	222 06	76 26	37 54	998771 989125 937028	987896	.7568	.6056	.7304
												Average ½Po .6327 qo .7919

No.	Qual.	Symbol	Letter	v	h	φ	ρ	$\lg \tan \varphi = \lg \frac{p_0}{q_0}$	$\frac{p_0}{q_0}$	$\frac{p_0}{q_0}$
11	P	$\infty 0$	a	347°13'	170°00'	-89°16'	90°00'	∞	1.608	1.608
12	F	∞	m	19 22	"	-58 07	"	020618	1.608	1.5908
13	G	$\infty 2$	g	38 58	"	-38 31	"	990061	.7954	1.5908
14	P	$\infty 4$	j	55 54	"	-21 35	"	959725	.3956	1.5814
15	P	$\infty 4$	j	99 35	"	22 06	"	960859	.4061	1.6244
16	F	$\infty 2$	g	115 56	"	38 27	"	989983	.7940	1.588
17	F	$\infty 0$	m	135 40	"	58 11	"	020731	1.612	1.612
18	P	$\infty 0$	a	167 27	"	89 58	"	∞	—	—
Average $\frac{p_0}{q_0} = 1.599$										

soon learned with a little practice. All the logarithms of col. 10 having been obtained, the numbers corresponding to each were found from the table, and entered in col. 11 opposite each. The upper number of each horizontal line in col. 11 is a value of $x = pp_0$; the lower a value of $y = qq_0$. p and q having been determined graphically (symbol, col. 3) the numbers of col. 11 yielded a series of values for p_0 ($\frac{1}{2}p_0$ in col. 12) and q_0 (col. 13) the average of which gave the elements of the crystal. The prisms yielded the ratio of p_0 to q_0 by a simple calculation as shown. The result of the calculation of this crystal were as follows:

$$p_0 = 1.2654; q_0 = 0.7919;$$

$$\frac{p_0}{q_0} = 1.597 \text{ (from prisms } \frac{p_0}{q_0} = 1.599).$$

A similar calculation may be made for each crystal measured and the results averaged. Possibly a simpler method is to average the angles for each form, make one calculation from these angles, and, weighting the resulting element values according to the number and quality of the readings for each form, find a final average. Table 2 shows the observed forms and angles (averaged) measured on eight crystals of higginsite together with the range of variation of each angle. The values of the elements p_0 and q_0 there given are the basis for the calculation of the φ and ρ angles of the same table. Table 3 shows the meth-

od of this calculation. The columns are numbered and the nature of the content of each is indicated by the heading. The lowest group of figures in each heading indicates the operation by which the values of the column were obtained. For example in col. 3, (1 + lg. p) means that lg. p is added to each logarithm of col. 1; in col. 5, (3-4) means that the logarithm of col. 4 is to be subtracted from that of col. 3; in col. 9 (3-6 = 4-7) means that the result of subtracting each logarithm of col. 6 from that of col. 3 should equal the result of subtracting each logarithm of col. 7 from that of col. 4. Incidentally this identity is a check on the calculation of each set of angles. After the values of lg. tan φ are obtained in col. 5, the angle is found from the table, entered in col. 8 and at the same time the lg. sin and lg. cos entered in cols. 6 and 7. ρ is obtained from lg. tan ρ of col. 9.

Each operation is thus a horizontal addition or subtraction or the result of adding a common value to a vertical series of numbers. All operations are supposed to be done mentally or with the aid of a slip of paper on which a commonly used value may be written. With comparatively little practice they become very easy and rapid.

In the Introduction to Goldschmidt's *Winkeltabellen* will be found forms similar to this for the calculation of crystals of each system. It was in this way that the enormous labor of calculating all the angles contained in that work was accomplished.

TABLE 2

HIGGINSITE. ANGLE TABLE OF CALCULATED AND OBSERVED VALUES.

$$p_0 = 1.272 \quad q_0 = .7940$$

	Symbol		Calculated		Measured		No. of Faces	Limits	
	Mill.	Gdt.	φ	ρ	φ	ρ		φ	ρ
a	100	$\infty 0$	90°00'	90°00'	90°00'	90°00'	12		
B	210	2∞	72 40	"	72 40	"	2	72°08'-73°12'	
C	320	2∞	67 24	"	68 03	"	4	67 20-68 48	
m	110	∞	58 02	"	58 03	"	7	57 48-58 15	
g	120	$\infty 2$	38 42	"	38 32	"	5	38 17-38 47	
J	140	$\infty 4$	21 50	"	21 52	"	9	20 59-22 33	
u	011	01	00 00	38 27	00 00	38 24	6		38°06'-38°47'
e	102	$\frac{3}{2} 0$	90 00	32 27	90 00	32 28	6		32 18-32 50
z	101	10	"	50 51	"	52 11	1		
y	302	$\frac{3}{2} 0$	"	62 20	"	62 20	3		62 04-62 39
o	111	1	58 02	56 18	58 02	56 09	7	57 39-58 18	55 54-56 30
p	122	$\frac{1}{2} 1$	38 42	45 30	38 38	45 38	7	38 00-39 40	44 50-46 00
r	342	$\frac{3}{2} 2$	50 14	68 04	50 15	68 06	7	50 00-50 39	67 19-68 35
s	528	$\frac{5}{8} 4$	75 59	39 20	76 41	39 25	2		
A	746	$\frac{7}{2} 6$	70 22	57 36	70 00	59 34	2		
x	326	$\frac{1}{2} 3$	67 24	34 34	67 53	34 33	2		

TABLE 3
TABLE TO SHOW METHOD OF CALCULATION OF ANGLES
(See *Winkeltabellen*, pp. 18, 19 & 19a).

Mineral Higgsinite Elements $p_0 = 1.272$ $\lg p_0 = 0.10449$ $q_0 = .7940$ $\lg q_0 = 989982$	Let. Symb. pq	1 $\lg p$	2 $\lg q$	3 $\lg x = \lg p p_0$ $1 + \lg p$	4 $\lg y = \lg q q_0$ $2 + \lg q$
	$\left[\begin{array}{c} o \ 1 \\ p \frac{1}{2} 1 \\ r \frac{3}{2} 2 \end{array} \right]$	0	0	010449	989982
		969897	0	980346	989982
		017609	030103	028058	020085
5	6	7	8	9	10
$\lg \frac{p p_0}{q q_0} = \lg \tan \varphi$ 3 - 4	$\lg \sin \varphi$ from 8	$\lg \cos \varphi$ from 8	φ from 5	$\lg \frac{p p_0}{\sin \varphi} = \lg \frac{q q_0}{\cos \varphi}$ $= \lg \tan \rho$ 3 - 6 = 4 - 7	ρ from 9
020467	992858	972381	58°02'	017591 017601	56°18'
990364	979605	989233	38 42	000741 000749	45 30
007973	988573	980595	50 14	039485 039490	68 04

LISTS OF THE ORTHORHOMBIC MINERALS INCLUDED IN GOLDSCHMIDT'S WINKELTABELLEN. EDGAR T. WHERRY. *Washington, D. C.*—As the prism zone is on the whole most characteristic of orthorhombic crystals, it has seemed desirable to arrange the minerals of this system in the order of increasing values of axis a .

ORTHORHOMBIC MINERALS

	a	c	Page		a	c	Page
Uranophanite	0.31	1.01	355	Topaz	0.53	0.95	346
Polycrasite (Poly- kras)	0.35	0.31	271	Pucherite	0.53	1.17	274
Euxenite	0.36	0.30	137	Phosphosiderite	0.53	0.88	266
Molybdate	0.39	0.47	243	Jordanite	0.54	1.02	191
Columbite	0.40	0.36	101	Yttrotantalite	0.54	1.13	371
Oanneroedite (An- nerödit)	0.40	0.36	45	Rammelsbergite	0.54	—	291
Flinkite	0.41	0.74	147	Samarските	0.55	0.52	309
Monticellite	0.43	0.58	253	Struvite	0.55	0.62	332
Fayalite	0.46	0.58	252	Mascagnite	0.56	0.73	232
Tephroite	0.46	0.59	254	Bertrandite	0.57	0.60	64
Hjelmite	0.46	1.03	177	Hopeite	0.57	0.47	180
Olivine	0.47	0.59	251	Beryllonite	0.57	0.55	66
Ardennite	0.47	0.31	53	Mica (Glimmer)	0.58	3.29	161
Chrysoberyl	0.47	0.58	97	Dyscrasite (Anti- monsilber)	0.58	0.67	49
Aeschynite	0.48	0.67	31	Argentopyrite (Silber kies)	0.58	0.55	318
Diaphorite	0.49	0.73	115	Stromeyerite	0.58	0.97	330
Pyrostilpnite (Feuer- blende)	0.50	0.70	145	Chalcocite (Kupfer- glanz)	0.58	0.97	205
Wavellite [old data]	0.50	0.38	362	Sternbergite	0.58	0.84	329

Caracolite	0.58	0.42	88	Langite	0.79	0.42	212
Iolite (Cordierit)	0.59	0.56	103	Daviesite	0.79	0.48	112
Niter (Kalisalpeter)	0.59	0.70	194	Hambergite	0.80	0.73	169
Bromlite (Alstonit)	0.59	0.74	34	Chalcostibite (Wolfs-			
Cotunnite	0.59	1.19	105	bergit)	0.80	0.63	367
Fischerite	0.59	—	147	Mendipite	0.80	—	237
Carnallite	0.60	1.39	88	Sulfur (Schwefel)	0.81	1.91	313
Frieseite	0.60	0.74	153	Barite (Baryt)	0.82	1.31	60
Thenardite	0.60	1.25	341	Bismite [trigonal?]	0.82	1.60	70
Orpiment (Auripig-				Jamesonite	0.82	—	187
ment)	0.60	0.67	57	Thermonatrite	0.83	0.81	341
Witherite	0.60	0.73	365	Pinakiolite	0.83	0.59	267
Euchroite	0.61	1.04	133	Haidingerite	0.84	0.99	168
Strontianite	0.61	1.04	331	Prehnite	0.84	1.12	272
Cerussite	0.61	0.72	89	Brookite	0.84	0.94	80
Schrockingerite	0.61	—	313	Manganite	0.84	0.54	230
Zoisite	0.62	0.34	379	Kornerupite	0.85	—	200
Aragonite	0.62	0.72	51	Serpierite	0.86	1.36	316
Stephanite (Melan-				Prismatite	0.86	0.83	273
glanz)	0.63	0.69	233	Mazapilite	0.86	0.99	233
Kentrolite	0.63	0.90	197	Strengite	0.87	0.98	330
Descloizite	0.64	0.80	113	Scorodite (Skorodit)	0.87	0.96	321
Polyhalite	0.64	—	270	Triphylite	0.87	1.05	350
Variscite	0.65	—	358	Enargite	0.87	0.82	127
Nesquehonite	0.65	0.46	248	Dufrenite (Kraurit)	0.87	0.43	201
Atacamite (Atakamit)	0.66	0.75	56	Pseudobrookite	0.87	0.89	274
Lawsonite	0.67	0.74	216	Nadorite	0.89	1.39	245
Ilvaite (Lievrit)	0.67	0.44	220	Zinkosite	0.89	1.41	374
Loellingite (Löllin-				Anhydrite	0.89	1.01	44
git)	0.67	1.23	223	Spodiosite	0.89	1.58	325
Lithargite (Bleioxyd)	0.67	0.98	72	Zinckenite	0.90	1.14	372
Sundtite	0.68	0.45	333	Ochrolite	0.91	2.01	250
Arsenopyrite (Arsen-				Hemafibrite (Häma-			
kies)	0.68	1.19	55	fibrit)	0.91	1.74	68
Glaucodotite (Glauc-				Reddingite	0.91	1.05	293
kodot)	0.69	1.19	160	Tellurite	0.92	0.93	339
Acanthite (Akanthit)	0.69	0.99	32	Caledonite	0.92	1.41	87
Erythrosiderite	0.69	0.72	132	Danburite	0.92	0.88	108
Staurolite	0.69	0.98	327	Goethite (Göthit)	0.92	0.61	162
Epigenite	0.69	—	131	Cosalite	0.92	1.46	104
Tungstite	0.70	1.61	352	Synadelphite	0.92	1.72	337
Hydrocyanite	0.71	1.26	186	Gerhardtite	0.92	1.16	156
Polymignite	0.71	0.51	271	Stilbite (Desmin)	0.93	0.76	113
Harstigit	0.71	1.01	171	Diasporite	0.94	0.60	116
Laurionit	0.73	0.83	214	Bournonite	0.94	0.90	76
Alloclasite (Alloklas)	0.74	0.55	34	Dufrenoyssite	0.94	1.53	120
Klaprothite	0.74	—	199	Krennerite	0.94	0.51	202
Marcasite (Marka-				Stylotypite	0.94	—	331
sit)	0.76	1.21	232	Meneghinite	0.95	0.69	238
Euchlorite	0.76	1.88	133	Olivenite	0.95	0.68	251
Fluellite	0.77	1.87	148	Lanthanite	0.95	0.90	213
Eosphorite	0.78	0.52	128	Uranothallite	0.95	0.78	355
Brochantite	0.78	0.49	393	Newberyite	0.95	0.94	249
Childrenite	0.78	0.53	93	Sartorite (Sklero-			
Celestite (Cölestin)	0.78	1.28	98	klas)	0.96	0.77	320
Calamine, Hemimor-				Libethenite	0.96	0.70	220
phite (Kieselzink-				Emplectite (Emple-			
erz)	0.78	0.48	197	ktit)	0.96	0.77	126
Lecontite	0.78	1.53	218	Bismuthinite (Wis-			
Valentinite	0.79	1.41	357	muthglanz)	0.97	0.99	364
Anglesite	0.79	1.29	42	Adamite	0.97	0.72	30

Patrinite.....	0.97	—	258	Thomsonite.....	0.99	1.01	342
Sillimanite.....	0.97	—	319	Leucophanite (Leu-			
Andorite.....	0.98	0.87	41	kophan).....	0.99	0.67	219
Goslarite (Zinkvit-				Ludwigite.....	0.99	—	224
riol).....	0.98	0.56	375	Cerite.....	1.0-	0.81	89
Natrolite.....	0.98	0.35	246	Uranospinite.....	1.0-	1.46	355
Morenosite (Nickel-				Geocronite (Geo-			
vitriol).....	0.98	0.57	249	kronite).....	1.01	0.58	156
Nagyagite.....	0.98	1.78	245	Enstatite [group]....	1.03	0.59	281
Gismondite.....	0.99	0.94	157	Hydromagnesite.....	1.04	0.47	186
Andalusite.....	0.99	0.70	40	Kermesite (Anti-			
Guarinite.....	0.99	0.74	166	monblende).....	1.32	0.85	46
Epsomite.....	0.99	0.57	132	Polybasite.....	1.73	1.58	270
Astrophyllite.....	0.99	4.70	55	Epididymite.....	1.74	1.85	128
Stibnite (Antimon-				Humite.....	2.20	1.08	181
glanz).....	0.99	1.02	47				

REPRESENTATIVES OF CLASSES WITH DIMINISHED SYMMETRY

CLASS HEMIMORPHIC

Struvite.....	0.55	0.62	Calamine, hemimorphite..	0.78	0.48
Bertrandite.....	0.57	0.60	Prehnite.....	0.84	1.12

CLASS SPHENOIDAL

Epsomite.....	0.99	0.57
Leucophanite.....	0.99	0.67
Edingtonite.....	1.0-	0.95

PERI-ORTHORHOMBIC

Mica group.....	Monoclinic
Polybasite.....	Monoclinic

NOTES AND NEWS

A CALCIUM PHOSPHATE WITH RATIOS BETWEEN THOSE OF TRIPLITE AND SARCOPSIDE. EDW. F. HOLDEN. *Hillsboro, N. H.*—In the writer's note on sarcopside in the May number of this magazine (pages 99-102), the formula-types of the various fluo-phosphates and related minerals were compared, in table 3; it was also noted in discussing that table that a ferrous fluophosphate from Stoneham, Maine, has been found to show a composition lying approximately midway between the sarcopside and apatite ratios, $R:(F, OH):(PO_4) = 12:3:7$. The purpose of the present note is to call attention to another apparently intermediate mineral, also from Stoneham, the analysis of which is given (as a peculiar "apatite") in *U. S. Geol. Survey Bull* 591, p. 349. The ratio derivable from this analysis is $11:6:4$, which is $\frac{2}{3}$ of the way from sarcopside to triplite ($7:2:4 + 2 \times (2:1:1) = 11:6:4$). The chief base in this mineral is calcium, so the member of the triplite group concerned is spodiosite; but the properties of the Stoneham mineral are so unlike those ascribed to spodiosite as to make its distinctness seem at least possible. Studies of the optical properties, with special reference to homogeneity, of minerals appearing to occupy intermediate positions in the series are necessary, however, before their status can be settled.

NOTE ON SULFUR AS A MINERAL OF THE MOON. EDGAR T. WHERRY, *Washington, D. C.*—While looking up information on the occurrence of the element sulfur, the writer came across a reference to its presence on the moon,¹ and as its identification there is probably not generally familiar to mineralogists, a note upon the matter has been prepared. On making a photograph of the moon by ultra-violet light, Professor R. W. Wood noticed the presence of a peculiar dark spot bordering the crater known as Aristarchus. A view of the same region taken thru a yellow screen showed no such spot, while one thru a violet screen showed it faintly. Trials of different terrestrial volcanic rocks with the same color screens showed that corresponding effects were obtained only in rocks containing a thin film of sulfur, no other known substance yielding exactly the same results. And, as the form and position of the Aristarchus spot suggests that it represents material thrown out from the crater by a volcanic blast, the conclusion is justified that it consists either of ash containing sulfur, or of a deposit of this element formed by condensation of ejected vapor. Here is a practically untouched field for research—developing of methods for identifying moon minerals—or should we say moonerals?

TWO AMERICAN OCCURRENCES OF EPIDESMINE. SAMUEL G. GORDON. *Academy of Natural Sciences of Philadelphia*—Epidesmine, the orthorhombic form of $(\text{Ca}, \text{Na}_2)\text{Al}_2\text{Si}_6\text{O}_{18} \cdot 6\text{H}_2\text{O}$ (the monoclinic form of which is represented by the common zeolite stilbite), was described by Rosicky and Thugutt² in 1913 from Schwarzenberg, where it occurred as a crust on calcite associated with orthoclase and fluorite. As no other localities have yet been reported,³ two American occurrences are worthy of note.

The mineral was collected by Mr. Frederick Oldach of the Reading High School, at a trap quarry one-half mile west of Robeson, or Gickerville, on the Schuylkill River, 7 miles south of Reading, Berks County. It occurs as small colorless or yellow prismatic crystals, a combination of the three pinacoids $a(100)$, $b(010)$, and $c(001)$; $a(100)$ is characteristically pearly, and sometimes slightly iridescent. The epidesmine is intimately associated with natrolite; other minerals noted in the quarry are stilbite (very abundant), prehnite, laumontite, chabazite, apophyllite, calcite, chrysocolla and epidote.

Col. Washington A. Roebling had previously identified the mineral at Moore Station, Mercer County, N. J. Optical examination of a specimen kindly presented by him to the writer showed it to be identical with the mineral from Schwarzenberg and Robeson.

The epidesmine from the three localities showed the following optical characters: optically —; $\alpha = 1.485$, $\beta = 1.495$, $\gamma = 1.500$, all $\pm .005$; $\gamma - \alpha = 0.015$. Axial plane parallel to $a(100)$; $a = Y$, $b = Z$, $c = X$; $Bx_a \perp$ to $c(001)$; $2E$ approximately 40° .

¹ Wood, R. W. Selective absorption of light on the moon's surface and lunar petrography. *Astrophys. J.*, 36, 75–84, 1912.

² V. Rosicky and St. J. Thugutt: Epidesmin, ein neuer Zeolith, *Centr. Min. Geol.*, 1913, 422–426; Ford: Third Appendix to Dana's System of Mineralogy, 27, 1915.

³ The "stilbite" figured by Heddle (Mineralogy of Scotland, II, plate LXXX, fig. 2, 1901), and by Böggild (Mineralogia Groenlandica, 562, fig. 108, 1905) is most probably epidesmine.

The John R. Stanton collection of mineral specimens has been sold to Mr. M. L. Morgenthau of New York City. The sale was made by G. S. Scott of New York who represented Mr. Stanton. The collection comprises 4500 specimens. For upwards of 30 years the collection was in the making, it having been originated by Mr. Stanton's brother and after his death Mr. John R. Stanton bought many very choice specimens and enlarged the collection considerably. Mr. Stanton's connection with the copper mines of Michigan secured for him what is probably the finest collection of native coppers in calcite known. Some of these calcites, clear as crystal and with native copper imbedded, make astonishingly beautiful specimens. Likewise Mr. Stanton's assortment of datolites from Michigan is unsurpassed, there being 100 or more beautiful polished specimens. His crystallized native silver and copper specimens from Lake Superior were many and choice. In crystallized copper Mr. Stanton had several splendid specimens, the finest being one weighing 35 kg. and showing hundreds of perfect crystals. There were many most attractive specimens in the collection representing various minerals from every prominent locality. Mr. Morgenthau is to be complimented in having secured this splendid mineral collection. The collection for years had been on exhibition at Mr. Stanton's office, 15 William St., New York City.

The Extension Division of the University of California is offering a correspondence course in determinative mineralogy, given by Professor Arthur S. Eakle. While intended primarily for residents of California, it is available for everyone, no matter where they live, and is open for enrollment at any time. Believing that this course might be of interest to many of our readers, we have obtained further details about it, which will be found on page i of this issue. For additional information address Professor Allyn G. Smith, Chairman, Technical Department, University of California, Berkeley, Cal.

Villamaninite

W. R. Schoeller and A. R. Powell: Villamaninite, a new mineral. *Min. Mag.*, 19, [88], 14-18, 1920.

NAME: After the village Villamanin, Cármenes district, Prov. Leon, Spain.

PHYSICAL PROPERTIES: Crystallization is cubic, with the octahedron and cubo-octahedron as the recognizable forms. $H. = 4\frac{1}{2}$. Color, iron black with a dull metallic luster. Streak, sooty black. No cleavage and uneven fracture. Sp. gr., 4.4-4.5.

CHEMICAL PROPERTIES: In closed tube gives sublimate of sulfur and selenium. Soluble in nitric acid with liberation of globule of sulfur. A sulfide of copper, nickel, cobalt, and iron, rich in selenium; probably a disulfide $(Cu, Ni, Co, Fe)(S, Se)_2$. Four analyses gave approx.: Cu 19, Ni 18, Co 7, Fe 4, S 50, Se $1\frac{1}{2}$ per cent. Traces were also found of arsenic, bismuth, lead and zinc, while negative tests are reported for tellurium, thallium, indium and gallium.

OCCURRENCE: Evenly disseminated thru a matrix of white crystalline dolomite, associated with chalcopyrite, iron pyrite and quartz. Occurs in groups of rough crystals and as small nodular masses with a radially fibrous structure.

W. F. HUNT.

[This material appears to be either a mixture, or a cupriferous polydymite. It is too poorly characterized to rank as a distinct and definite species.]

W. F. F.]